



The global distribution and environmental drivers of aboveground versus belowground plant biomass

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A poor understanding of the fraction of global plant biomass occurring belowground as roots limits our understanding of present and future ecosystem function and carbon pools. Here we create a database of root-mass fractions (RMFs), an index of plant below- versus aboveground biomass distributions, and generate quantitative, spatially explicit global maps of RMFs in trees, shrubs and grasses. Our analyses reveal large gradients in RMFs both across and within vegetation types that can be attributed to resource availability. High RMFs occur in cold and dry ecosystems, while low RMFs dominate in warm and wet regions. Across all vegetation types, the directional effect of temperature on RMFs depends on water availability, suggesting feedbacks between heat, water and nutrient supply. By integrating our RMF maps with existing aboveground plant biomass information, we estimate that in forests, shrublands and grasslands, respectively, 22%, 47% and 67% of plant biomass exists belowground, with a total global belowground fraction of 24% (20–28%), that is, 113 (90–135) Gt carbon. By documenting the environmental correlates of root biomass allocation, our results can inform model projections of global vegetation dynamics under current and future climate scenarios.

The extent and distribution of carbon storage in plants is central to our understanding of the structure and function of the terrestrial biosphere^{1,2}. At present, our understanding of global plant biomass stems primarily from satellite and inventory data that characterize the aboveground components of ecosystems^{1,3–6}. However, a major component of plant biomass cannot be detected by satellites or direct aboveground observations because it exists belowground as roots, which are central to the functioning of terrestrial ecosystems⁷ and greatly affect ecosystem nutrient, water and carbon cycling^{8–11}. Plant investment in root biomass has generally been explored using local-scale observations^{12,13}, such that variation at regional and global scales remains unknown. Without empirical, spatially explicit models of the relative proportion of plant biomass distributed to roots (root-mass fraction, RMF), aboveground biomass cannot be adequately scaled to total plant biomass^{5,14,15}.

At the global scale, patterns of biomass fractions in roots, wood and leaves are expected to reflect a functional balance between the distribution of resources that exist above (light, CO₂) and below (water, nutrient) the soil surface^{16,17}. Indeed, plants tend to show proportionally higher biomass allocation to roots when belowground resources are scarce^{16–20}. Because belowground limitations should be most pronounced in cold regions (where nutrient supply is low) and in dry environments^{13,21–23}, RMFs are expected to increase with latitude²⁴ and in arid regions with low water availability^{13,22}. Given such relationships, changing climate and associated shifts in above- versus belowground resource availability may alter relative plant investments in root biomass. Due to large variations in the longevity of different plant tissues, such shifts drive feedbacks between climate and the terrestrial carbon cycle²⁵. Yet, despite the central role of biomass allocation in shaping vegetation responses to altered environmental conditions, we lack a quantitative, spatially explicit understanding of the factors driving RMF^{13,19,26}, which is

essential for improving confidence in future carbon cycle projections²⁵. Characterizing the environmental drivers of the variation in RMF across the globe is thus critical for projecting plant biomass accumulation²³, productivity²⁷ and nutrient cycling¹¹ under current or future climate change scenarios⁶.

Here we use a large dataset of plant RMFs to generate a spatially explicit understanding of the proportional plant biomass that exists in roots and upscale these predictions to the globe to quantify total belowground plant biomass. First, we constructed a database of 17,814 plot-level RMF measurements, aggregated to 6,803 spatially distinct samples from forest ($n=5,170$), grassland ($n=1,293$), and shrubland ($n=340$) vegetation (Fig. 1). Using 63 global layers of climate, topography, soil and vegetation features (Extended Data Fig. 1), we then trained random forest models to characterize the environmental factors that drive geographic variations in RMFs within each vegetation type and interpolate these relationships across the globe. To comprehend the interactive mechanisms that have ultimately shaped variations in biomass partitioning within and across regions, we additionally use recursive partitioning and regression methods accounting for spatial autocorrelation. Finally, we compare our empirical RMF data with vegetation model simulations to pinpoint important deviations in model predictions versus observed biomass fractions.

Global mapping of RMFs and model performance

There was considerable variation in RMF within each vegetation type (Fig. 2a–c), which can rival the amount of variation observed across types. Within forests, the highest RMFs are found in boreal regions and seasonally dry tropical and subtropical regions, such as the Mediterranean region, east Brazil, California, and Southern Africa (Fig. 3a). Within grasslands, RMFs in tropical areas are, on average, ~50% lower than in temperate environments, and the

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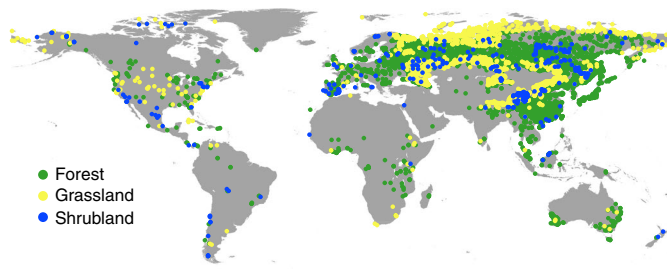


Fig. 1 | RMF sample locations in forests, shrublands and grasslands. A total of 17,814 RMF records (13,620 forest, 2,468 grassland and 1,726 shrubland) were collected and aggregated into 6,893 spatially unique data points used for geospatial modelling (5,170 forest, 1,293 grassland and 340 shrubland records; Methods).

highest RMFs are found in montane and temperate steppe regions with low temperature and soil moisture, such as the Qinghai-Tibetan and Mongolian plateaus (Fig. 3c). In shrublands, RMFs are highest in the boreal and arid biomes (Fig. 3e), while low RMFs occur in flooded savannas and tropical and subtropical ecosystems.

Based on a spatial validation procedure (spatially buffered leave-one-out cross-validation, LOO-CV), our random forest models explained 16–39% of the spatial variation in RMFs within vegetation types (LOO-CV $R^2=0.16$, 0.34 and 0.39 for forests, shrublands and grasslands, respectively; Fig. 2d–f; Extended Data Fig. 2a–c shows random tenfold cross-validation results). We used LOO-CV to account for the potential effects of spatial autocorrelation on model validation statistics²⁸, with buffer radii of 250 km in forests and shrublands and 500 km in grasslands. These buffer ranges were based on the extent of autocorrelation in the model residuals (Extended Data Fig. 2g–i shows residual semivariograms and Fig. 2g–i shows LOO-CV R^2 s for buffer radii from 1 to 500 km).

We then used these predictive relationships to spatially upscale observed RMFs over existing forest, shrubland and grassland areas across the globe to reveal the global distribution of RMFs for each vegetation type and combined across all existing vegetation (excluding cropland) (Fig. 3). Given the considerable variability in these pixel-level RMF estimates, the predictive uncertainty is relatively high at the single-pixel level. As such, although our map is projected at 30 arcsec (~ 1 km² at the Equator) resolution, high-precision predictions require distances larger than 10 km (Extended Data Fig. 2d–f). Yet, it is worth noting that, even at these distances, systemic bias could potentially cause these predictions to deviate from the true mean. We attempted to minimize this possibility by collecting data from a wide number of sources that would be unlikely to represent the same systematic bias, but the risk remains, particularly in areas with low sampling density. For example, due to the low sampling density (Fig. 1), our models had high prediction uncertainty in African grasslands (Fig. 3d).

The environmental controls of RMFs

To generate a mechanistic understanding of the environmental factors driving the global variation in RMFs, we explored the directional effects of eight environmental features that have previously been shown to influence RMFs, including annual mean temperature and water availability (annual mean soil moisture and aridity index)^{21–23} (Fig. 4 and Extended Data Figs. 3–6). Our analyses show that, across all vegetation types, temperature and water availability are the dominant drivers of RMFs at the global scale (Fig. 4a–c), supporting the idea that belowground biomass increases when belowground resources are scarce^{18,19,29,30}. However, we also identify important interactions among these main drivers that can obscure the simple directional effects at a global scale (Figs. 4d–l and 5).

In forests, previous meta-analyses²¹ and intraspecific comparisons^{13,31} showed an effect of soil water availability on RMFs. However, a recent global study on forest biomass distribution did not detect an effect of water availability and instead indicated mean annual temperature as the main driver of RMF variation²³. Our global analyses can now reconcile this discrepancy, revealing strong interactions between temperature and soil moisture (Figs. 4d,g,j and 5 and Extended Data Fig. 4a,b) that probably result from geographic variation in the interplay between water, temperature and nutrient limitation^{19,29,32}. In dry areas, temperature exhibited a positive effect on RMFs (Figs. 4d,j and 5a and Extended Data Fig. 4a), probably because the extent of water limitation there is exacerbated by hot temperatures, leading to an increased investment into building roots. In contrast, in wet areas, nutrient limitation should be the driving force of above- versus belowground biomass investment, explaining why we see increasing RMFs toward cold areas where nutrients are at low supply and are less mobile. The normalized difference vegetation index (NDVI), an indicator of photosynthetic activity and canopy cover³³, was negatively correlated with RMFs (Extended Data Fig. 3), suggesting reduced allocation to roots in dense forests where aboveground competition for light is high.

Shrublands follow similar trends to forests. The geographic gradients of RMFs are systematically affected by the interplay between temperature and water availability (Fig. 4e,h and Extended Data Fig. 5). As in forests, we find that the negative effect of temperature on RMFs strongly increases toward moist regions (Fig. 4e and Extended Data Fig. 5a). In grasslands, soil moisture and temperature also represent dominant controls of RMFs. In contrast to forests and shrublands, the negative effect of soil moisture on RMFs is most pronounced at intermediate temperature levels and the negative effect of temperature is most pronounced at dry locations (Fig. 4f,i and Extended Data Fig. 6), probably reflecting functional differences between woody and herbaceous species (grasses), as grasses might be more tolerant to low temperature and drought.

Despite the general impacts of water availability and temperature on RMFs, soil characteristics, such as sand content, depth to bedrock and soil organic carbon-to-nitrogen (C:N) ratio, also play an important role in the geographic RMF gradient across all vegetation types (Extended Data Figs. 3–6). For instance, RMFs in forests are higher in regions with sandy and deep soil (Extended Data Figs. 3 and 4d). High RMFs in sandy soils can probably be explained by reduced soil moisture due to the low water-holding capacity and an increased need for root stability, while lack of space belowground probably explains why shallow soils correlate with low RMFs. The C:N ratio of soil organic matter, a proxy for plant N availability^{34,35}, was positively correlated with RMFs in shrublands, but showed no correlation in forests and a negative correlation with RMFs in grasslands (Extended Data Fig. 3). Despite usually being nutrient poor, arid regions typically have very low soil C:N ratios as a result of a small organic C pool³⁶, explaining the negative correlation between soil C:N ratio and RMFs in grasslands and emphasizing the need for a direct characterization of nitrogen availability to represent global variations in plant-available nutrients.

Comparison with existing biomass models

To examine the possible utility of these spatial models, we explored the extent to which the observed variation in forest root-mass fractions is reflected in existing biomass models. We assessed the ability of four dynamic global vegetation models that contribute to the dynamic global vegetation model project TRENDY (version 8)³⁷ as well as the latest static global biomass prediction⁶ to replicate the observed RMFs in forest systems. The selected dynamic vegetation models do not currently simulate the spatial variation in RMFs that is reflected in our data (Extended Data Fig. 7; all R^2 values < 0). The most recent static model of global above- and belowground plant carbon⁶ also failed to explain the observed spatial variation in RMFs

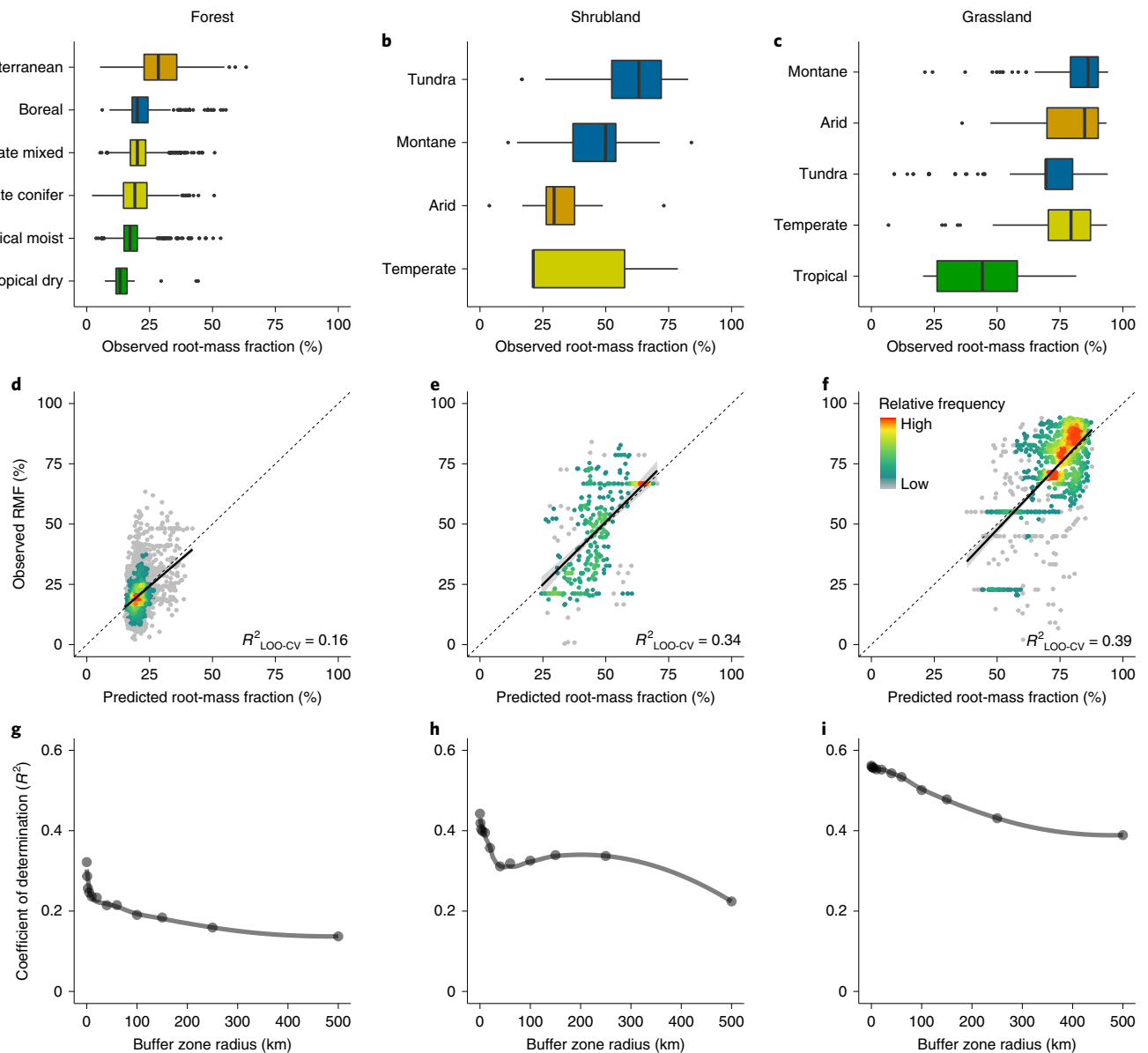


Fig. 2 | RMF variation and model validation. **a–c**, Boxplots of observed RMFs per vegetation type and biome ($n=5,170$ forest **(a)**, 340 shrubland **(b)** and 1,293 grassland **(c)** observations). **d–f**, Heat plots showing the relationships between predicted and observed RMFs in forests **(d)**, shrublands **(e)**, and grasslands **(f)** based on spatially buffered LOO-CV. Solid lines indicate fitted relationships based on ordinary least squares regression (coefficient-of-determination values relative to the 1:1 line [equation (2)] shown in the bottom right corner); dashed diagonal lines indicate a 1:1 relationship between observed and predicted points. **g–i**, LOO-CV coefficient-of-determination values (equation (2)) for buffer radii for data exclusion from 1 to 500 km in forests **(g)**, shrublands **(h)** and grasslands **(i)**. For each data point, all data within the respective buffer radius were excluded and all sample points out of the buffer zone were used to train the model.

(Extended Data Fig. 7b, $R^2 < 0$) and overpredicted RMFs by, on average, ~9.8%. As such, consideration of these new, emerging patterns in belowground biomass may be useful for enhancing the performance of global dynamic vegetation models. By capturing some of this unexplained variation in RMF across the globe, we hope that our predictive models can be used to benchmark and parameterize process-based models to represent the meaningful variation in plant belowground biomass investment across time and space.

Discussion

By reflecting the morphological adaptations of plants in different ecosystem types to varying resource availability, our global models of RMF are directly relevant for building a thorough understanding

of the plant carbon sink, and associated plant trait variation over time and space. For instance, roots often have a lower litter quality (high C:N ratio) relative to leaves and aboveground wood^{38,39}. Various studies additionally point towards a trade-off between plant growth rate and RMF, with high RMFs being associated with lower growth rates and higher stress tolerance^{38,40}. Given the significance of RMFs for ecosystem functioning on multiple levels^{7,27}, our study provides a baseline for future research allowing for upscaling these feedbacks to the global scale. Moreover, this study represents a critical step toward assessing the impacts of climate change on terrestrial carbon distributions. The strong interactive effects of temperature and water availability on RMFs observed here indicate that warmer and drier climates in the future might

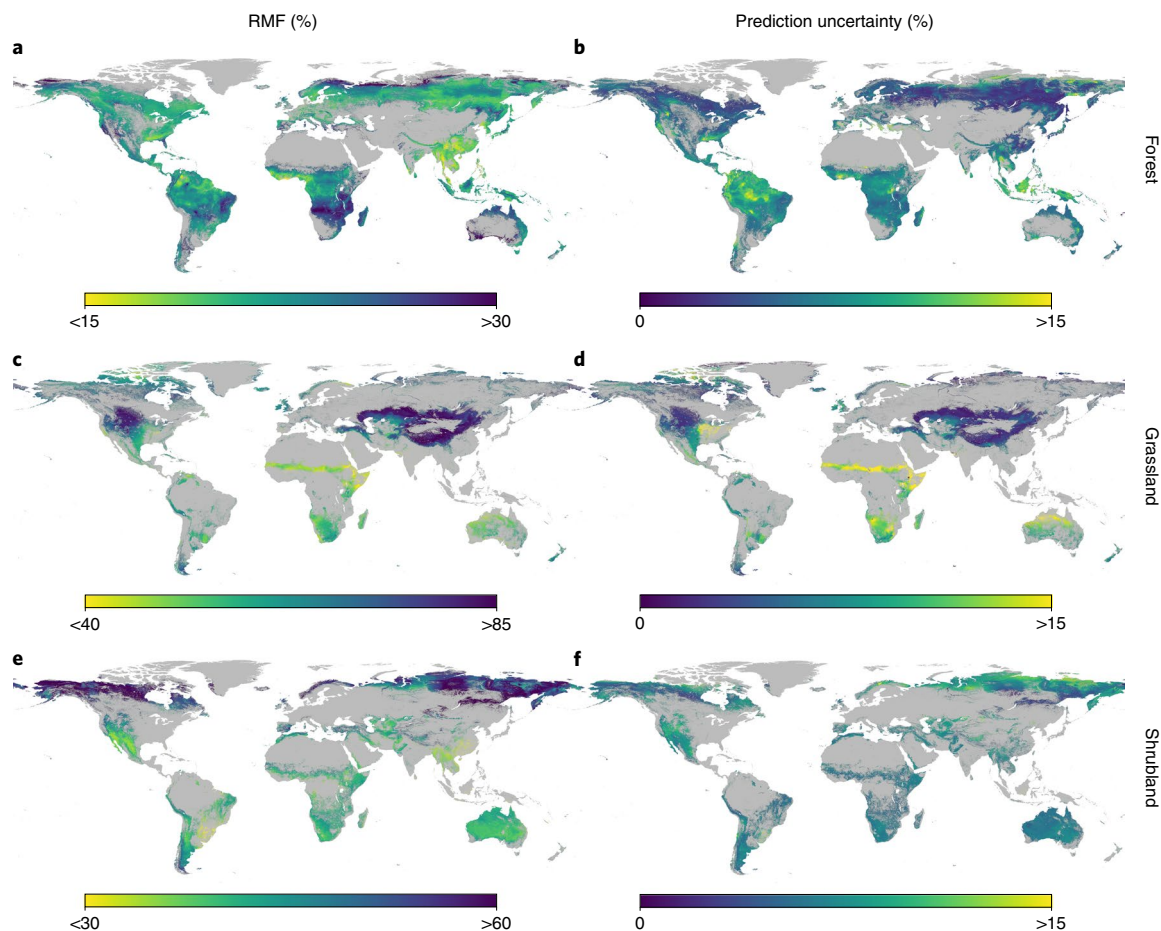


Fig. 3 | The global distribution of RMFs in forests, grasslands and shrublands. a,c,e, Predicted RMFs from random forest models based on empirical data from 5,170 forest (a), 1,293 grassland (c) and 340 shrubland (e) records. **b,d,f,** Bootstrapped variation coefficients (standard deviation divided by mean) in forests (b), grasslands (d) and shrublands (f). In forests, high uncertainty occurs in Southeastern Asia and central Amazonia. In grasslands, high uncertainty is found in African savannah and steppes, where we lack sampling data. All maps are projected at 30 arcsec ($\sim 1 \text{ km}^2$) resolution.

reshape the relative biomass allocation to root systems to maintain the functional balance of aboveground versus belowground resources and sinks.

The systematic variation in global plant RMFs that we detect now offers the opportunity to derive quantitative, spatially explicit estimates of total belowground biomass from existing aboveground biomass products. To exemplify this, we combined existing maps of aboveground plant biomass with land-cover maps⁴¹ and our vegetation-type-specific RMF estimates to approximate the variation in root biomass across the globe¹⁵ (Fig. 6 and Extended Data Fig. 8). We estimate a total belowground biomass of 93.9 (minimum to maximum = 77.5–113.8), 11.5 (9.1–14.4) and 5.2 (3.7–7.0) GtC for forests, shrublands and grasslands, respectively. As such, the contribution of belowground biomass to the total terrestrial plant biomass ranges around 24% (20–28%), highlighting the importance of belowground plant carbon pools for global biogeochemical cycling. This estimation of the total proportion of belowground biomass largely agrees with previous estimates that ranged from 20% to 30% at the global scale^{6,15,26,29,42,43}.

While our model predictions well reflect the observed variation in RMFs across large spatial gradients, there are still limits regarding our ability to accurately predict at finer spatial scales. First, although our observations support much of the mechanistic and experimental research in plant biomass investment, we must highlight that the correlations observed here are not necessarily indicative of causal relationships. In addition,

a considerable proportion of the variation in root allocation remains unexplained, underscoring the need for further research and data collection to improve predictions of plant biomass partitioning. All RMF uncertainty estimates in this study reflect the bootstrap prediction uncertainty (Fig. 3b,d,f) or the intermodel consistency (Extended Data Fig. 9) and thus cannot represent other potentially important sources of uncertainty, such as measurement error and sampling bias. Identifying ways to incorporate these other sources of uncertainty is a key outstanding challenge and will be critical to improving future modelling efforts. Some of the unexplained variation in our models is probably due to missing meta-information about the sites where the RMFs were observed. Because plot-level metadata is neither available nor could it be used for global extrapolation, our models had to rely on interpolated global layers of environmental conditions. These layers might accurately describe plot-level variations in aboveground parameters, such as mean annual temperature, but are probably introducing considerable uncertainties with respect to soil information, in particular nutrient content. Yet, despite these covariate uncertainties, the out-of-sample model fit of 16–39% (obtained from spatial LOO-CV; Fig. 2d–i) demonstrates the important role that climate and soil characteristics play in shaping plant biomass distributions above- and belowground.

In conclusion, our analysis reveals the spatial distribution of plant root allocation at a global scale that can inform our understanding of vegetation carbon storage and dynamics.

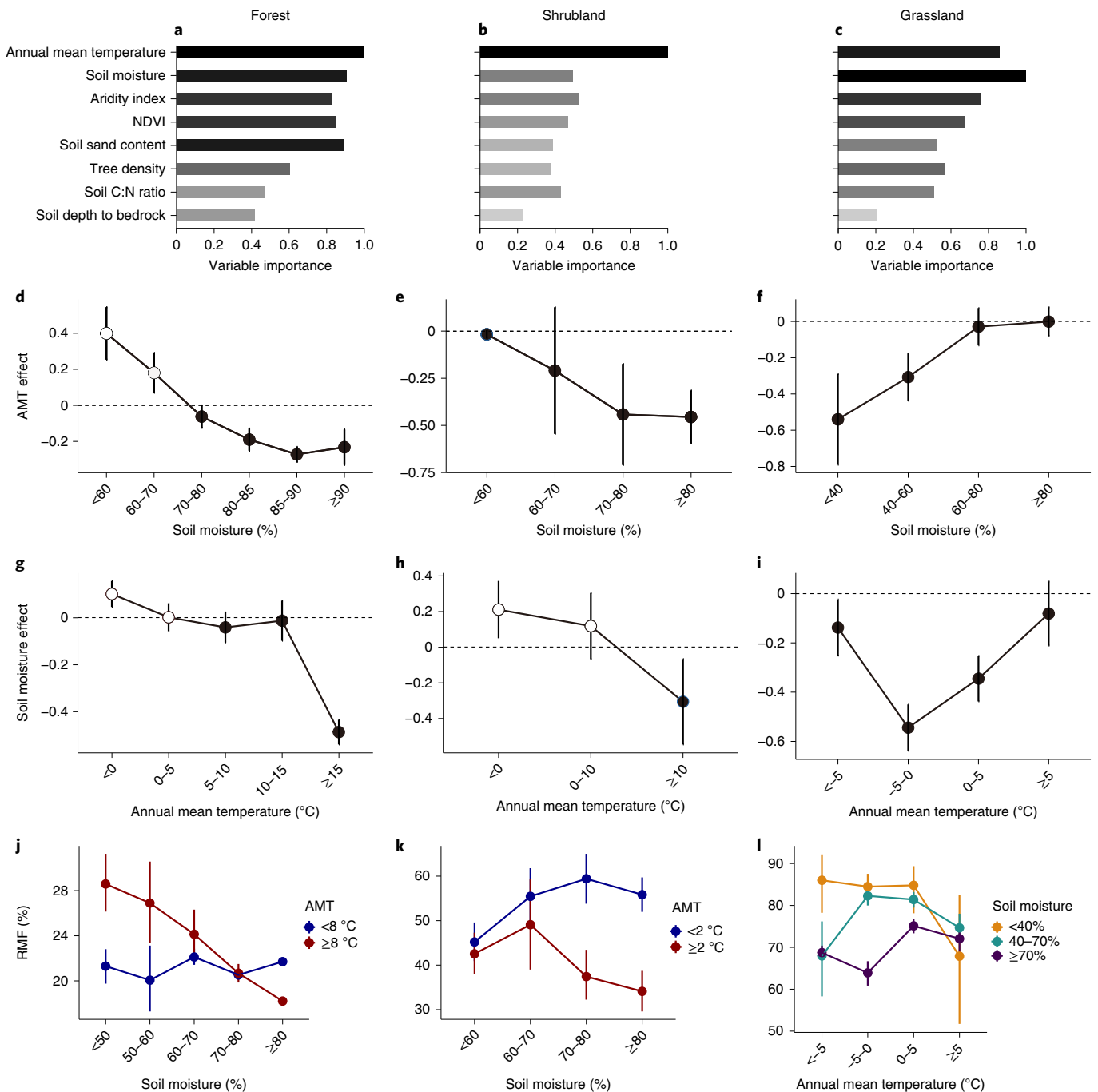


Fig. 4 | Relationships between environmental variables and RMFs. a–c, Variable importance metrics from random forest models for the effects of eight environmental variables on RMFs in forests (**a**), shrublands (**b**) and grasslands (**c**). **d–i,** The effects of annual mean temperature (AMT) (**d–f**) and soil moisture (**g–i**) on RMFs for different soil-moisture or temperature classes, respectively. Values represent mean standardized coefficients (± 2 standard errors) bounded between -1 and 1 , where negative values indicate a negative effect (closed circles) and positive values a positive effect on RMFs (open circles). **j–l,** Illustration of changes in mean RMFs ($\pm 95\%$ confidence intervals) along temperature and soil-moisture gradients. For forests (**j**) and shrublands (**k**), RMFs are shown in response to soil moisture for two temperature classes (red = high temperature, blue = low temperature); for grasslands (**l**), RMFs are shown in response to annual mean temperature for three soil-moisture classes (purple = high soil moisture, teal = intermediate and orange = low).

The geographic gradients in above- versus belowground biomass distribution contribute to our understanding of limiting factors to plant growth and offer a new fine-scale framework to estimate the total belowground carbon stocks of plants. We estimate that approximately one quarter of the total plant carbon is stored belowground, stressing the importance of root systems for the terrestrial carbon sink. These estimates of root biomass are directly

determined by existing estimates of aboveground biomass, which are characterized by considerable uncertainty. As we continue to expand our understanding of global aboveground biomass, these predictive models of proportional belowground biomass contributions will be valuable for improving our predictions of total plant biomass and ecosystem functioning under current and future climate scenarios.

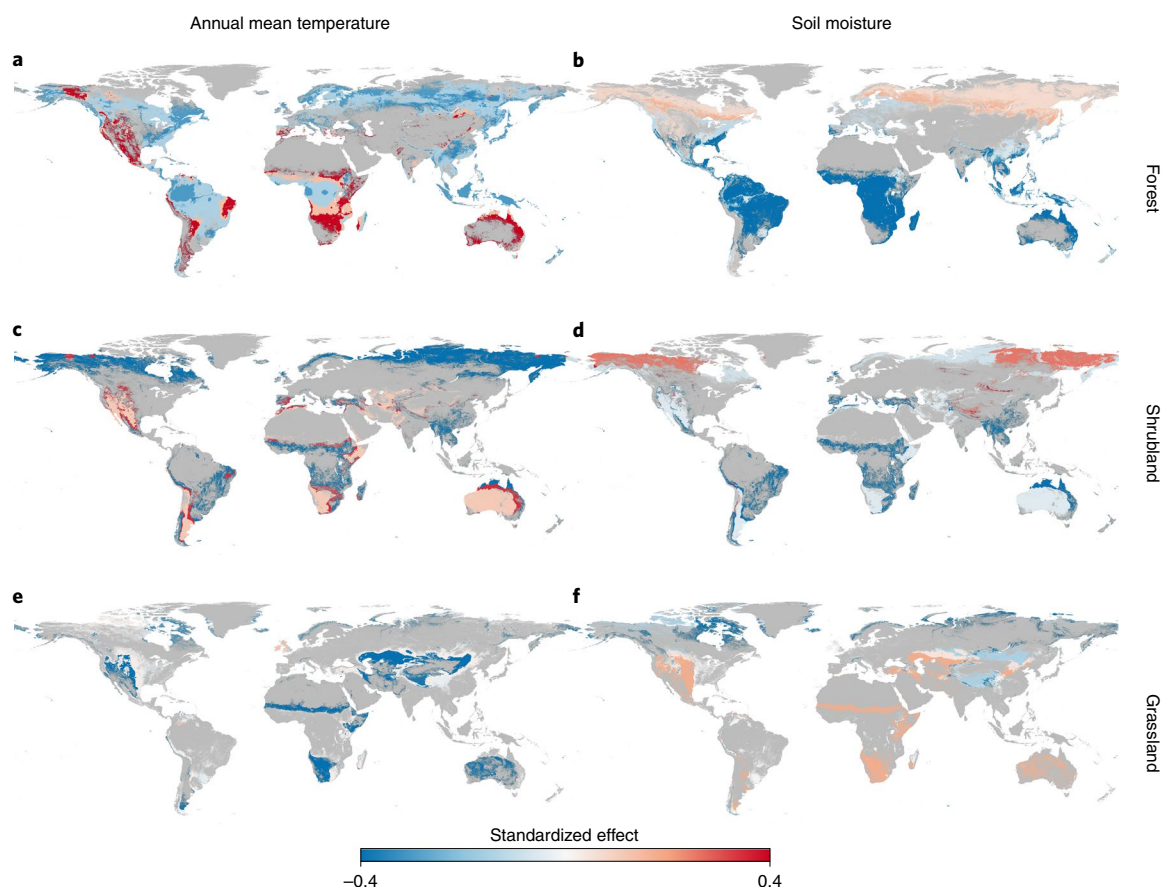


Fig. 5 | Spatial variations in the environmental correlates of RMFs. a–f, Maps show the relationships between RMFs and annual mean temperature (left panels) or soil moisture (right panels) in forests (**a,b**), shrublands (**c,d**) and grasslands (**e,f**). Effect sizes (standardized coefficients) for each region are based on recursive partitioning trees (Extended Data Figs. 4–6). Red colours represent a positive effect of annual mean temperature or soil moisture on RMFs and blue colours represent a negative effect.

Methods

Data collection. *Collection of plot-level RMFs.* To collect data on the biomass ratio between roots and shoots (root–shoot ratio, RSR) from the literature, we set ubiquitous criteria for all three vegetation types (forests, shrublands and grasslands) and specific criteria for some of the vegetation types. Data was only used if all of the following four criteria were met: (1) RSR values were available at the plot level (and not only for selected individuals); (2) data came from field measurements (allometry-derived RSRs were excluded). If the study consisted of multiple treatments, only data from the control treatment was selected; (3) coordinate information was given; and (4) measurements were taken after 1960. For datasets that included information on more than one vegetation type, the abovementioned criteria were evaluated separately for the subset of trees, grasses and shrubs within the dataset. Furthermore, to minimize the influence of ontogenetic drift on RSRs in forests (higher root-to-shoot biomass ratios in seedlings than in adults⁴⁴), we (1) excluded studies that focused on tree seedlings and (2) excluded seedlings or tree individuals within studies with a diameter at breast height (DBH) smaller than 10 cm. This selection process resulted in a total of 17,814 plot-level RSR records, with 13,620 forest records from refs.^{12,21,45–95}, 2,468 grassland records from refs.^{26,63,81,82,96–113} and 1,726 shrubland records from refs.^{58,63,81,82,91,96,99,105,114–184}. For each vegetation type, we aggregated records from the same location by calculating the mean and removed obvious outliers by visual inspection (0.7%, 2.0% and 4.9% of the forest, grassland and shrubland data, respectively), resulting in spatially distinct data for a total of 5,170 forest, 1,293 grassland and 340 shrubland records (Fig. 1). These aggregated measurements thus reflect the ratio of living plant biomass existing below- versus aboveground in a given location. We then converted all root–shoot ratio records to root-mass fractions (RMFs), restricting the data to values from 0 to 100%:

$$\text{RMF} = \frac{\text{RSR}}{\text{RSR} + 1} \times 100\%. \quad (1)$$

Environmental covariates. In total, 63 environmental layers, including climate, soil, anthropogenic, vegetation and topographic variables, were used as covariates in

our analyses (Extended Data Fig. 1 and ref.¹⁸⁵ provide covariate layer details). All layers were standardized to 30 arcsec resolution (~1 km² at the Equator).

Geospatial modelling of root-mass fractions. Random forest models. To train global models of RMFs for each vegetation type, we ran a series of random forest machine learning models using the R package *h2o*¹⁸⁶. We used a grid-search procedure to iteratively explore the results of a suite of random forest models trained on the 63 covariates. A list of hyperparameters was selected, varying in (1) the number of trees (50, 100 or 150), (2) the number of variables sampled at each split (2 to 10) and (3) the minimum observations per leaf (2 to 5), resulting in a total of 108 models for each vegetation type. During the grid-search procedure, we assessed the performance of each model using random tenfold cross validation. Coefficient-of-determination values (R^2) for each fold were used to compute mean and standard deviation values for the cross-validated model. We selected the best ten models for each vegetation type based on the highest mean R^2 values.

All R^2 values reported throughout the manuscript represent the coefficient of determination relative to the 1:1 line of observed versus predicted values. This is equivalent to a standardized mean squared error:

$$R^2 = 1 - \frac{\text{RSS}}{\text{TSS}} = 1 - \frac{n \times \text{MSE}}{\text{TSS}} = 1 - \frac{\sum_{i=1}^n (O_i - P_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (2)$$

where RSS refers to the residual sum of squares, TSS the total sum of squares, MSE the mean squared error, O_i the observations, P_i the predictions and n the number of observations. This metric reflects systematic deviation between model predictions and observed values. Negative values indicate that the model performs worse than using the mean across observations.

To create the final maps, we used an ensemble approach, whereby, for each vegetation type, we averaged the global predictions from the best ten random forest models. The final R^2 values (based on the ensemble of the cross-validated predictions) for RMFs in forests, shrublands and grasslands were 0.33, 0.47 and 0.58, respectively (Extended Data Fig. 2a–c). By taking the average prediction across multiple models, ensemble methods minimize the influence of any single

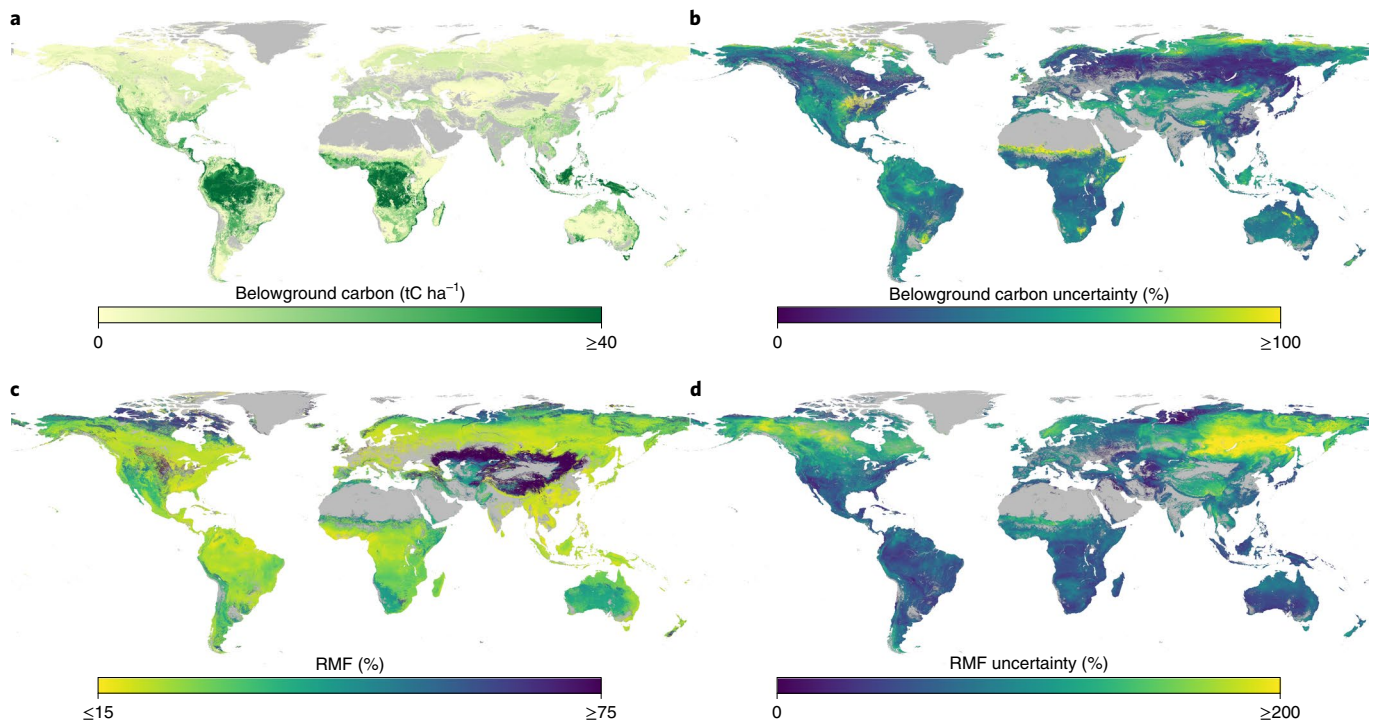


Fig. 6 | The global distribution of belowground plant biomass. **a**, Average belowground plant biomass per area (tonnes of carbon per hectare) scaled by the relative biomass contribution of each vegetation type (Methods). **b**, Relative prediction uncertainty calculated as the difference between the highest and lowest belowground carbon estimates per pixel (based on the minimum and maximum estimates for each vegetation type) divided by the mean carbon estimate in the respective pixel. **c,d**, Average pixel-level RMF distribution (**c**) and associated uncertainty (**d**) integrating forest, grassland and shrubland RMFs. To obtain pixel-level averages, RMFs of the three functional types were weighted by the relative contribution of each type to the aboveground biomass in the respective pixel (Methods). With this map, pixel-level estimates of aboveground biomass can be directly scaled to belowground biomass. The uncertainty map shows the difference between maximum and minimum RMFs divided by averaged RMFs. All maps are projected at 30 arcsec ($\sim 1 \text{ km}^2$) resolution.

prediction, thereby stabilizing variation and minimizing bias that can otherwise arise from extrapolation or in-fit overfitting when using a single machine learning model¹⁸⁷. Moreover, by quantifying the variation across these ensemble predictions (intermodel consistency), we can identify areas that have low consensus across multiple models (Extended Data Fig. 9).

To implement this ensemble approach, the mean predicted value across the ten best-fitting models was used as the final model prediction for each vegetation type for each pixel, and the coefficient of variation across these ten models was used to characterize intermodel consistency (see “Evaluating model uncertainties”). All predictions were masked by the global distribution of forests, shrublands and grasslands ($\geq 10\%$ coverage), based on a 30 arcsec (1 km^2) consensus land-cover map⁴¹, representing the percentage of area covered by each vegetation type per pixel.

We also generated a unified RMF map, reflecting a weighted RMF average across vegetation types (Fig. 6c). To do so, we accounted for the relative aboveground biomass contribution of each vegetation type to the pixel-level total biomass (see “Belowground biomass estimation”).

To test for the potential effect of spatial autocorrelation in the model residuals, which can cause overestimation of model predictive power²⁸, we ran spatially buffered LOO-CV for a series of buffer radii from 1 to 500 km. In this procedure, each data point is predicted by a model that uses all data outside the buffer radius of the respective data point as training data, resulting in 5,170, 1,293 and 340 independent model runs for forests, grasslands and shrublands, respectively. For each vegetation type, we used the hyperparameter settings of the best-performing random forest model (based on random tenfold cross validation, see above). As in the other models, model performance was evaluated by the coefficient of determination relative to the 1:1 line (equation (2)) of observed versus predicted values (Fig. 2d–i).

To assess the range of autocorrelation, we created semivariograms of model residuals (Extended Data Fig. 2g–i). The residual semivariograms of our random cross-validation and leave-one-out cross-validation models show autocorrelation ranges on the order of ~ 150 to 500 km. Based on these results, we chose buffer radii of 250 km for forests and shrublands and 500 km for grasslands, which represent the range of autocorrelation of RMF residuals.

All models were run in R version 3.5.1 (ref. 188). Geospatial mapping was performed in Google Earth Engine¹⁸⁹.

Evaluating model uncertainties. To represent the confidence of our estimates for each pixel based on sampling uncertainty, we performed a stratified bootstrapping procedure using all data within each vegetation type. We set the number of bootstrap iterations to 100, that is, all pixels of the raster layer data (93,693,855 in forests, 60,977,011 in shrublands and 49,464,610 in grasslands) were classified 100 times. Each of the three datasets (forest, shrubland and grassland) were sampled with replacement, using biome as stratification category to proportionally represent the major bioclimatic zones in each of the 100 bootstrap samples. The size of each bootstrap sample was equal to the size of the original datasets. Each of the 100 bootstrap samples were then used to classify the composite image. Finally, these 100 images were used to calculate per-pixel coefficient-of-variation values (standard deviation divided by the mean predicted value) as a measure of sampling uncertainty (hereafter referred to as bootstrap prediction uncertainty; Fig. 3b,d,f).

To identify pixel-level uncertainties in the unified RMF map, combining all three vegetation types, we first calculated the difference between the upper (97.5% quantile) and lower (2.5% quantile) RMF prediction for each pixel and vegetation type, based on the bootstrapping procedure. We then obtained the lower and upper RMFs within each pixel, based on the lower and upper RMFs in each vegetation type, accounting for the relative aboveground biomass contribution of each vegetation type to the pixel-level total biomass (Fig. 6d and see “Belowground biomass estimation”).

We additionally created a spatially explicit visualization of the variation in predictions across the ten best random forest models, hereafter referred to as intermodel consistency, by calculating the per-pixel coefficient of variation (standard deviation divided by the average) across the ten best random forest models (Extended Data Fig. 9).

To assess the extent of extrapolation in our models, that is, how well our training data represents the full multivariate environmental covariate space, we performed a principal-component-analysis–based approach following van den Hoogen et al.¹⁸⁵. We projected the covariates composite into the same space using the centring values, scaling values and eigenvectors from the principal component analysis of the training data. Then, we created convex hulls for each of the bivariate combinations from the top principal components (which collectively covered more than 90% of the sample space variation), and classified whether each pixel falls within or outside each of these convex hulls. We used 17 principal components

(PCs) with 136 combinations for all covariates for the forest dataset, 15 PCs with 105 combinations for all covariates for the shrubland dataset, and 16 PCs with 120 combination for all covariates for the grassland dataset. This analysis revealed that 94%, 70% and 78% of land pixels falling in the extent of the respective vegetation type are covering at least 90% of the environmental conditions present in our training data locations in forest, shrubland and grassland, respectively (Extended Data Fig. 10).

The maps are available online at: <https://code.earthengine.google.com/2b17d6844f6077ea25021988cbc9b4c7>.

Environmental drivers of geographic variation in root-mass fractions.

Variable selection. To explore the effect of magnitude and direction of the main environmental drivers of RMFs within forests, shrublands and grasslands, we selected 8 uncorrelated environmental factors among our 63 covariates, representing climate, vegetation and soil features that are of high ecological importance and that have shown significant relationships with RMF in previous studies^{21–23,82}: annual mean temperature, aridity index, NDVI, tree density, soil organic C:N ratio, soil depth to bedrock, soil-moisture profile (the available water for the plant divided by the total water-holding capacity of the soil) and soil sand content. To test for multicollinearity among covariates, we calculated variance inflation factors using the R package *HH*¹⁹⁰. For all vegetation types, variance inflation factors were less than three, indicating sufficient independence among predictor variables.

Variable importance testing. We then ran a random forest model, implemented in the *h2o* R package¹⁸⁶, and quantified the relative importance of each of the eight variables, as determined by calculating whether that variable was selected to split on during the tree-building process, and how much the squared error (over all trees) improved as a result. To test whether our subset of 8 variables is representative of the main drivers of RMFs across our full random forest model including all 63 covariates, we compared the R^2 values (from random tenfold cross validation) of the reduced 8-covariate model with the full 63-covariate model. The results show that the reduced model captures 78%, 79% and 83% of the variation explained by our full model for forests, grasslands and shrublands, respectively, indicating that our subset of 8 independent variables represents the main drivers of our full model.

To further explore how the effects of environmental variables on RMFs vary depending on the magnitude of other environmental drivers, we ran recursive partitioning analyses using the packages *party*¹⁹¹ and *ggparty*¹⁹². Using a machine learning (that is, decision tree) algorithm, we explored the context dependency of the slope and intercept of a univariate linear model for the effect of any of the top four predictor variables (Fig. 4a–c) on RMFs, while the three remaining variables were evaluated as potential split points (Extended Data Figs. 4–6 and Fig. 5). The minimum node size (minimum number of observations contained in each terminal node) was set to ~10% of the total data, that is, 500 for forests, 120 for grasslands, and 30 for shrublands. These analyses demonstrate that the effects of one variable are highly dependent on other variables; the slope of the regression even changed from negative to positive depending on other explanatory variables (for example, Extended Data Fig. 4a). To exemplify the strong interactive effects of temperature and water availability, we additionally characterized the effects of mean annual temperature on RMFs within different soil-moisture classes and vice versa (Fig. 4d–l).

Partial-regression analysis. We also tested for consistent global trends of the eight variables on RMFs by univariate partial-linear-regression models. For forests, we ran bootstrapped partial-linear-regression models to remove the spatial autocorrelation present in our full dataset (see below) by randomly subsampling 1,700 points (33%) with replacement out of a total of 5,170 points 1,000 times. For each subsample, we fit a partial-regression model for each covariate of interest and calculated the corresponding slope. The results were then aggregated by calculating the mean coefficient (slope) across all model runs. For grasslands and shrublands, we used the whole dataset for our univariate partial-linear models. These partial-regression models show the overall effect of one variable when controlling for the global variation in all other variables. We tested for spatial autocorrelation in the residuals that could potentially bias significance tests and parameter estimates¹⁹³ by performing Moran's I tests. We detected significant spatial autocorrelation in the full forest data ($I=0.27$, $P<0.001$), and our bootstrapping procedure successfully removed this spatial autocorrelation (Moran's I P value >0.05) for 98% of the subsampled datasets. There was no significant spatial autocorrelation in grasslands ($I=0.04$, $P=0.36$) and shrublands ($I=0.15$, $P=0.13$).

As illustrated by the strong interaction between soil moisture and annual mean temperature (Fig. 4), a non-significant effect in these global linear models (Extended Data Fig. 3) does not mean that the respective variable does not affect RMFs; it rather indicates that there is no globally coherent trend detectable for this variable. For example, in shrublands, soil moisture had a positive effect on RMFs in cold regions and a negative effect on RMFs in warm regions (Fig. 4h). This dependency on other variables therefore explains why our global partial-regression models show no consistent general trend for this variable (Extended Data Fig. 3).

Belowground biomass estimation. Belowground biomass was estimated by integrating global estimates of aboveground biomass (Intergovernmental Panel on Climate Change (IPCC) tier 1)¹⁵ with our high-resolution maps of RMFs. To account for the occurrence of different vegetation types within pixels, we divided the aboveground biomass of each pixel into the relative proportion comprising different vegetation forms, derived from a high-resolution land-cover map⁴¹. Generally, total aboveground biomass of each pixel (AGB_i) can be partitioned by the following formula:

$$AGB_i = AGB_f + AGB_g + AGB_s \quad (3)$$

where AGB_f , AGB_g and AGB_s refer to the aboveground biomass of forests, grasses and shrubs for each pixel. For each vegetation type, the aboveground biomass can be rewritten as the product of aboveground biomass density (ABD; biomass per unit area), focal vegetation coverage (C), and pixel area (A_i). As such, aboveground biomass for each vegetation type can be expressed as:

$$AGB_f = ABD_f \times C_f \times A_i \quad (4)$$

$$AGB_g = ABD_g \times C_g \times A_i \quad (5)$$

$$AGB_s = ABD_s \times C_s \times A_i \quad (6)$$

and we can therefore rewrite equation (3) to:

$$AGB_i = (ABD_f \times C_f + ABD_g \times C_g + ABD_s \times C_s) \times A_i \quad (7)$$

Because the IPCC tier 1 biomass layer does not account for different vegetation types within pixels¹⁵, we estimated the biomass that each vegetation type contributes within each pixel by using fixed biomass ratios per unit area for each ecoregion zone¹⁵:

$$\alpha = \frac{ABD_f}{ABD_g} \quad (8)$$

$$\beta = \frac{ABD_s}{ABD_g} \quad (9)$$

where α is the biomass ratio between forests and grasslands and β is the biomass ratio between shrublands and grasslands. To obtain the biomass ratios α and β , we used aboveground biomass values for each vegetation type and ecoregion zone published in the IPCC report¹⁵. By combining these ratios with the spatial coverage of each vegetation type within pixels⁴¹, we were able to allocate the total aboveground biomass of each pixel known from equation (7) to the three vegetation forms as:

$$AGB_i = (C_f \times \alpha \times ABD_g + ABD_g \times C_g + C_s \times \beta \times ABD_g) \times A_i \quad (10)$$

where A_i is the area of the pixel and C_g , C_f and C_s refer to the relative grassland, forest and shrubland coverage.

Because all parameters in equation (10) are known besides ABD_g , we can solve this equation for each pixel to estimate ABD_g . We then used this value to obtain the corresponding aboveground densities for forest and shrubland using equations (8) and (9). Finally, we used these density values to obtain the total aboveground biomass for each vegetation type for each pixel, using equations (3–6)^{13,15}.

Using these pixel-level estimates of aboveground biomass for each vegetation type, we then calculated belowground biomass (BGB) per pixel by scaling these values by our RMF estimates using equation (1) (Extended Data Fig. 8a,c,e).

$$BGB_f = AGB_f \times RSR_f \quad (11)$$

$$BGB_g = AGB_g \times RSR_g \quad (12)$$

$$BGB_s = AGB_s \times RSR_s \quad (13)$$

BGB uncertainty within vegetation types was based on the upper and lower biomass estimates in each pixel and vegetation type, which were calculated from the upper and lower RMFs based on the bootstrapping procedure (Extended Data Fig. 8b,d,f).

To compute the total BGB in each pixel (Fig. 6a), BGB values of each vegetation type were aggregated via:

$$BGB_i = BGB_f + BGB_g + BGB_s \quad (14)$$

Total BGB uncertainty in each pixel was then derived from the sum of the upper or lower biomass estimates for each vegetation type (Fig. 6b).

To compute total belowground biomass at the global scale, we summed pixel-level BGBs:

$$\text{Total BGB}_i = \sum \text{BGB}_{i_r} \quad (15)$$

Similarly, belowground biomass of each vegetation type was summed by:

$$\text{Total BGB}_f = \sum \text{BGB}_{f_r} \quad (16)$$

$$\text{Total BGB}_g = \sum \text{BGB}_{g_r} \quad (17)$$

$$\text{Total BGB}_s = \sum \text{BGB}_{s_r} \quad (18)$$

To obtain a pixel-level RMF map that accounts for all three vegetation types (Fig. 6c), the belowground biomass of each pixel was divided by the total (below- and aboveground) biomass of each pixel:

$$\text{RMF}_i = \frac{\text{BGB}_i}{\text{AGB}_i + \text{BGB}_i} \quad (19)$$

Comparison of observed RMFs with dynamic global vegetation model outputs. We derived root-mass fractions as root biomass divided by total vegetation biomass and averaged (mean) over the years 1986–2015 from outputs of TRENDY (version 8) global vegetation model simulations following the S2 setup (that is, forced with observed climate and CO₂, while land use and land-use change were held constant at year 1700 state). The resulting maps represent ‘ecosystem-level’ root-mass fractions, aggregated across all plant functional types (PFT) by grid cell. To represent forest RMFs, we used only simulated RMFs of grid cells with a forest-cover fraction (determined by model outputs of total fractional plant coverage of all tree PFTs) of >50% (Extended Data Fig. 7). We tested the robustness of model–data comparison using an alternative threshold of >90% forest cover. Both thresholds yielded the same results, with R²s of observed versus predicted RMFs consistently <0. Models were selected for our analysis if outputs were provided for all required variables (root biomass, total vegetation biomass and PFT cover fraction).

Reporting Summary. Further information on research design is available in the Nature Research Reporting Summary linked to this article.

Data availability

The root–shoot ratio data underlying this study are available at https://github.com/haozhima95/Global_mapping_root_shoot_ratio/tree/master/RSR_data. Citations for the root–shoot ratio data are provided in the methods.

Code availability

The code used for this study is available at https://github.com/haozhima95/Global_mapping_root_shoot_ratio.git.

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Author contributions

H.M., L.M., T.W.C. and C.M.Z. conceived and developed the study and wrote the manuscript. H.M. and L.M. collected the data. H.M. and L.M. performed the analyses. D.S.M., J.v.d.H., B.S. and C.T. gave input on the manuscript.

Competing interests

The authors declare no competing interests.

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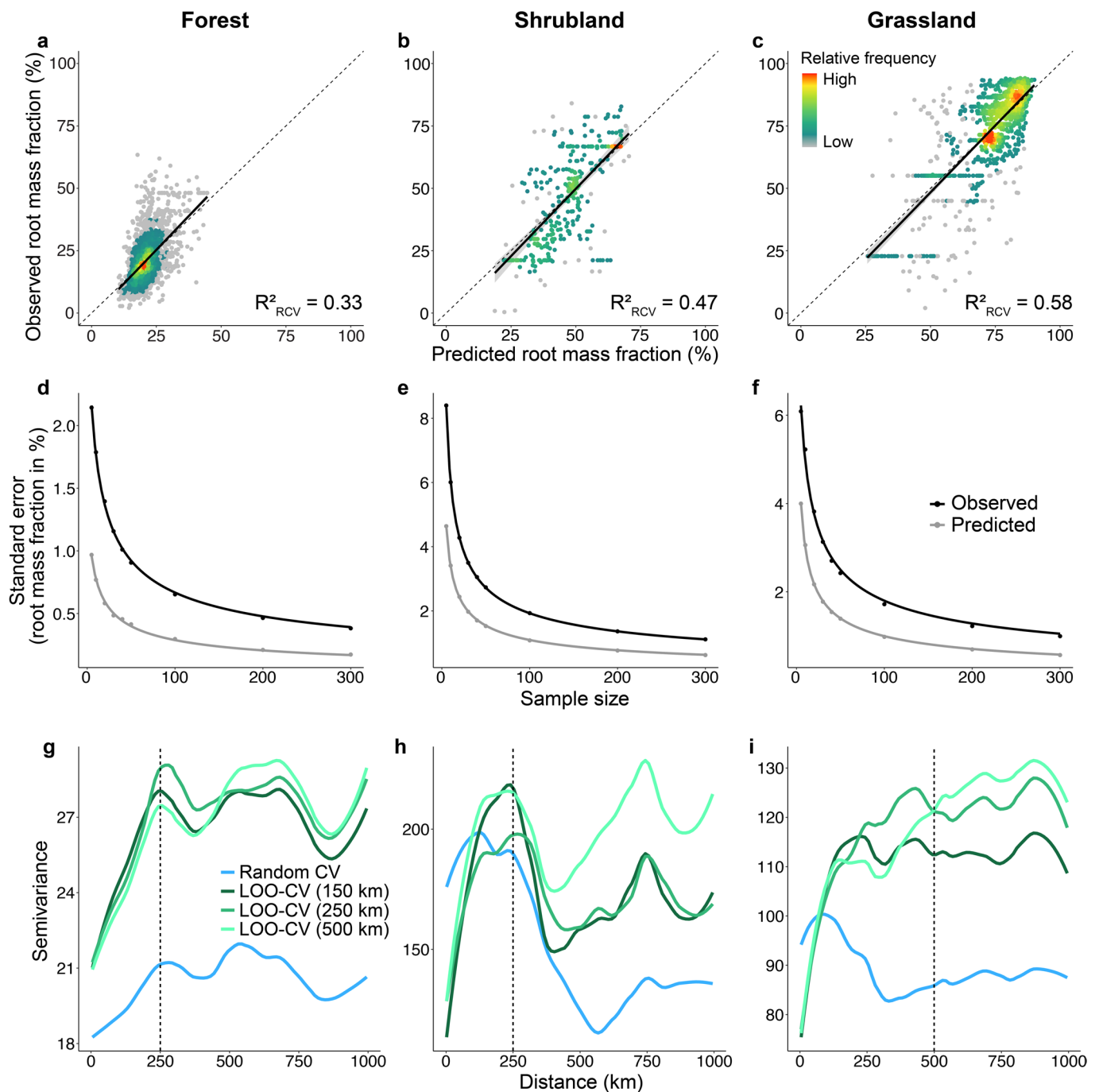
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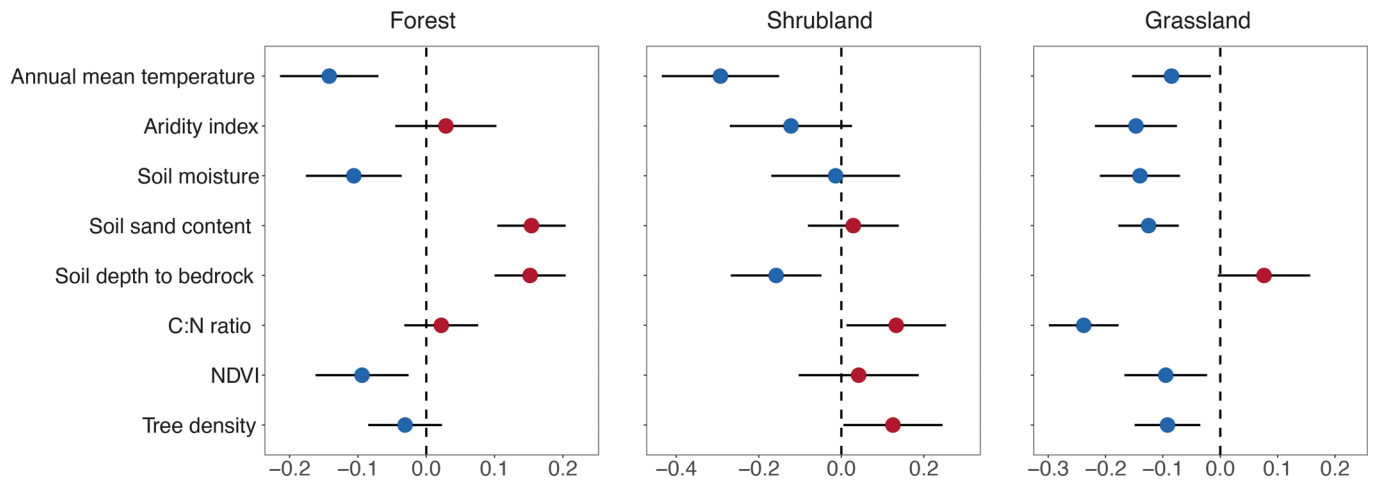
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Variable	Variable group	Original spatial resolution	Units	Source	
Annual mean radiation	Atmospheric	≈1km	W/m2	ref ¹⁹⁴	
Highest weekly radiation	Atmospheric	≈1km	W/m2		
Lowest weekly radiation	Atmospheric	≈1km	W/m2		
Radiation of coldest quarter	Atmospheric	≈1km	W/m2		
Radiation of driest quarter	Atmospheric	≈1km	W/m2		
Radiation of warmest quarter	Atmospheric	≈1km	W/m2		
Radiation of wettest quarter	Atmospheric	≈1km	W/m2		
Radiation seasonality	Atmospheric	≈1km	W/m2		
Annual mean temperature	Climatic	≈1km	°C	ref ¹⁹⁵	
Annual precipitation	Climatic	≈1km	mm		
Isothermality	Climatic	≈1km	%		
Maximum temperature of warmest month	Climatic	≈1km	°C		
Mean diurnal range	Climatic	≈1km	°C		
Mean temperature of coldest quarter	Climatic	≈1km	°C		
Mean temperature of driest quarter	Climatic	≈1km	°C		
Mean temperature of warmest quarter	Climatic	≈1km	°C		
Mean temperature of wettest quarter	Climatic	≈1km	°C		
Minimum temperature of coldest month	Climatic	≈1km	°C		
Precipitation of coldest quarter	Climatic	≈1km	mm		
Precipitation of driest month	Climatic	≈1km	mm		
Precipitation of driest quarter	Climatic	≈1km	mm		
Precipitation of warmest quarter	Climatic	≈1km	mm		
Precipitation of wettest month	Climatic	≈1km	mm		
Precipitation of wettest quarter	Climatic	≈1km	mm		
Precipitation seasonality	Climatic	≈1km	mm		
Temperature annual range	Climatic	≈1km	°C		
Temperature Seasonality	Climatic	≈1km	°C		
Global aridity index	Climatic	≈1km	Unitless		ref ¹⁹⁶
Global potential evapotranspiration (PET)	Climatic	≈1km	PET value		
Global population density	Anthropogenic	≈1km	persons/km2	ref ¹⁹⁷	
Human development percentage	Anthropogenic	≈1km	%	ref ¹⁹¹	
Human footprint	Anthropogenic	≈1km	Unitless	ref ¹⁹⁸	
Bulk density at depth 0~15cm	Soil	≈250m	kg/m3	ref ¹⁹⁹	
Cation exchange capacity of soil at depth 0~15cm	Soil	≈250m	cmole/kg		
Clay content at depth 0~15cm	Soil	≈250m	%		
Coarse fragments volumetric at depth 0~15cm	Soil	≈250m	%		
Depth to bedrock	Soil	≈250m	cm		
Probability of occurrence in the R horizon	Soil	≈250m	%		
Sand content at depth 0~15cm	Soil	≈250m	Mass fraction in %		
Silt content at depth 0~15cm	Soil	≈250m	Mass fraction in %		
Soil organic carbon content at depth 0~15cm	Soil	≈250m	g/kg		
Soil organic carbon stock at depth 0~15cm	Soil	≈250m	tones/hectare		
Soil pH*10 in H2O at depth 0~15cm	Soil	≈250m	pH*10		
Soil moisture profile	Soil	≈1km	%		ref ²⁰⁰
Soil water table depth	Soil	≈1km	m		ref ²⁰¹
Soil Organic C:N ratio	Soil	≈1km		ref ²⁰²	
Soil total Nitrogen	Soil	≈1km	g/kg		
Eastness	Topographic	≈1km	-	Global Multi-resolution Terrain Elevation Data 2010	
Elevation	Topographic	≈1km	Meters		
Hillshade	Topographic	≈1km	Grayscale value		
Northness	Topographic	≈1km	-		
Slope	Topographic	≈1km	Degree		
BRDF Albedo mandatory quality band 1	Vegetative	≈500m	reflectance band value	ref ²⁰³	
BRDF Albedo mandatory quality band 2	Vegetative	≈500m	reflectance band value		
BRDF Albedo mandatory quality band 3	Vegetative	≈500m	Reflectance band value		
BRDF Albedo mandatory quality band 4	Vegetative	≈500m	reflectance band value		
BRDF Albedo mandatory quality band 5	Vegetative	≈500m	reflectance band value		
BRDF Albedo mandatory quality band 6	Vegetative	≈500m	reflectance band value		
BRDF Albedo mandatory quality band 7	Vegetative	≈500m	reflectance band value		
Global 16-day enhanced vegetation indices (EVI)	Vegetative	≈250m	EVI value	ref ²⁰⁴	
Global 16-day vegetation indices (NDVI)	Vegetative	≈250m	NDVI value		
Tree density	Vegetative	≈1km	Number of trees/ha	ref ²⁰⁵	

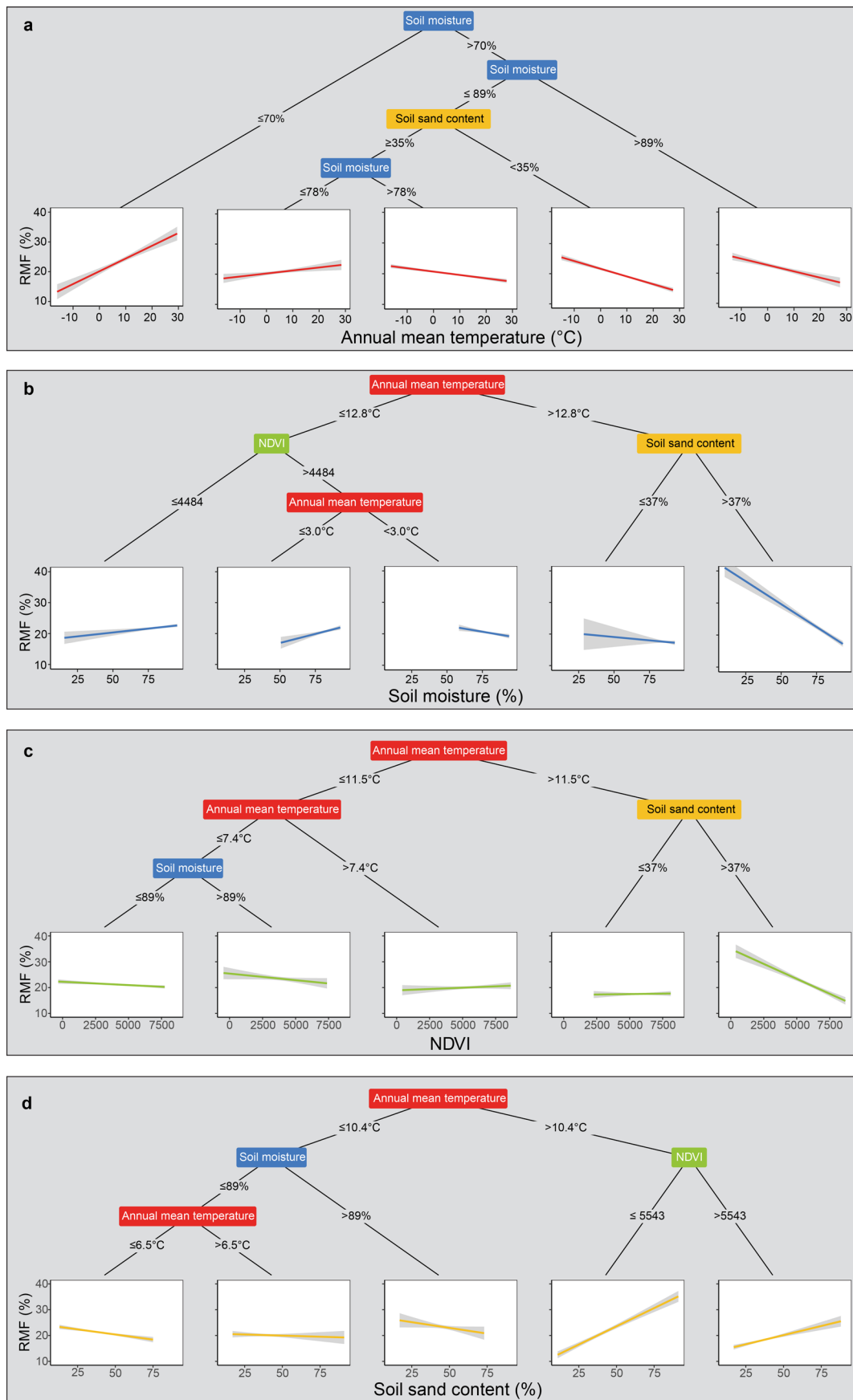
Extended Data Fig. 1 | Information on the 63 selected covariate layers used to model root mass fractions¹⁹⁴⁻²⁰⁵.



Extended Data Fig. 2 | Random 10-fold cross-validation (RCV) of the spatial root-mass fraction models and spatial autocorrelation of model residuals. **a-c**, Heat plots showing the relationships between predicted and observed RMFs in forests (**a**), shrublands (**b**), and grasslands (**c**) based on RCV. Solid lines indicate fitted relationships based on ordinary least squares regression [coefficient of determination values relative to the 1:1 line (equation 2) shown in the bottom right corner], dashed diagonal lines indicate a 1:1 relationship between observed and predicted points. **d-f**, The standard errors of the observed (black) and predicted (grey) mean values of root mass fractions decrease with increasing sample size. The operation was repeated with 1,000 random seeds for the observed and predicted mean values, and the calculated standard errors of the mean are shown. Note, 'sample size' in D-F refers to the number of pixels, and thus denotes square kilometres. **g-i**, Semivariograms illustrating spatial autocorrelation of model residuals in forests (**g**), shrublands (**h**) and grasslands (**i**). Semivariances of residuals were computed based on random 10-fold cross validation (blue) and spatial leave-one-out cross-validation (LOO-CV) with buffer radii of 150km (dark green), 250km (green) and 500km (light green). Dashed vertical lines indicate the buffer radii of the final validation model reported throughout the text.

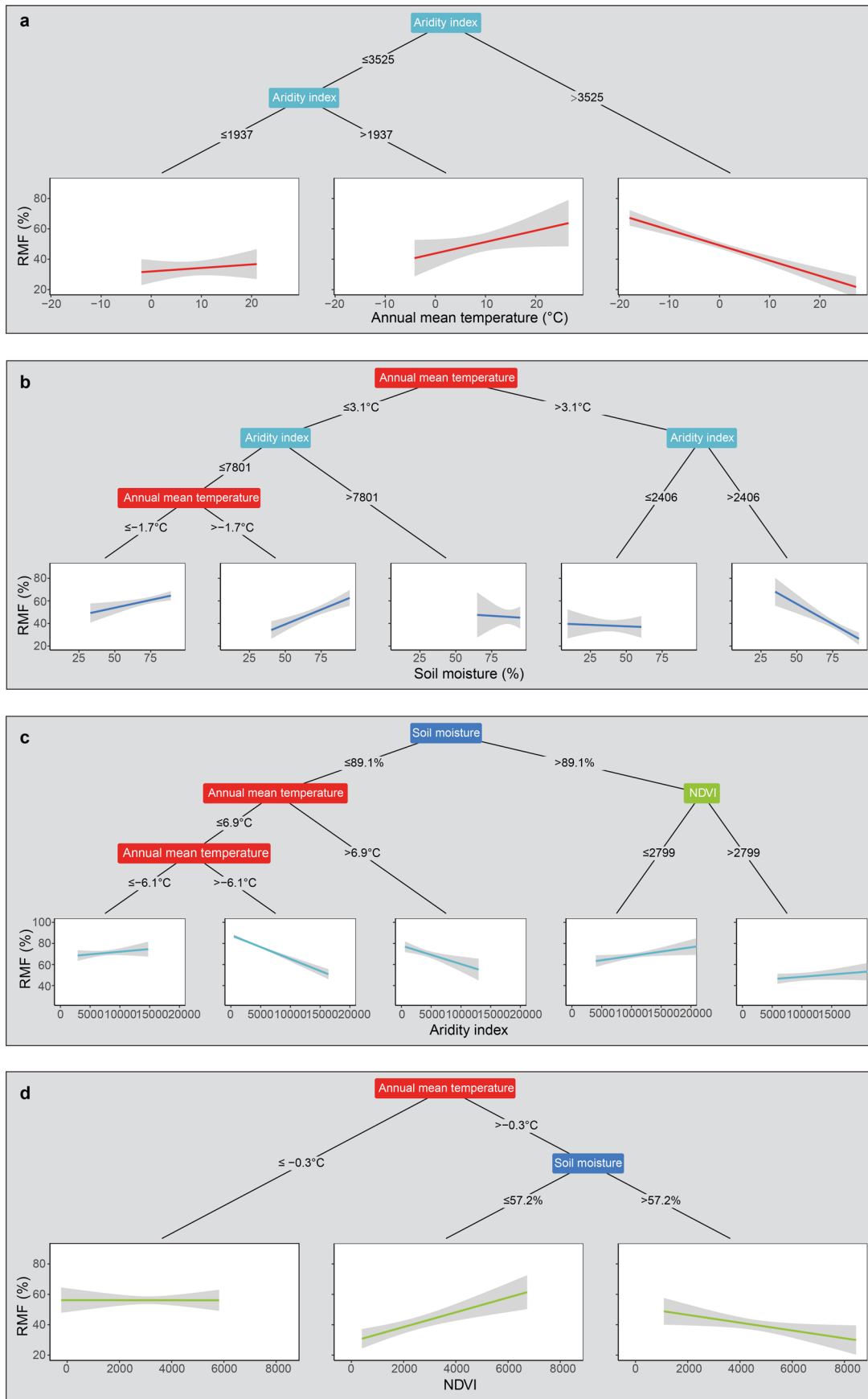


Extended Data Fig. 3 | Partial regression coefficients for the effects of 8 environmental covariates from linear multiple regression models. To reduce the influence of spatial autocorrelation, a bootstrapping procedure was applied for the forest data (see Methods). Red dots indicate positive effects on RMFs, blue dots indicate negative effects. Error bars reflect two standard errors either side of the mean partial regression coefficient.



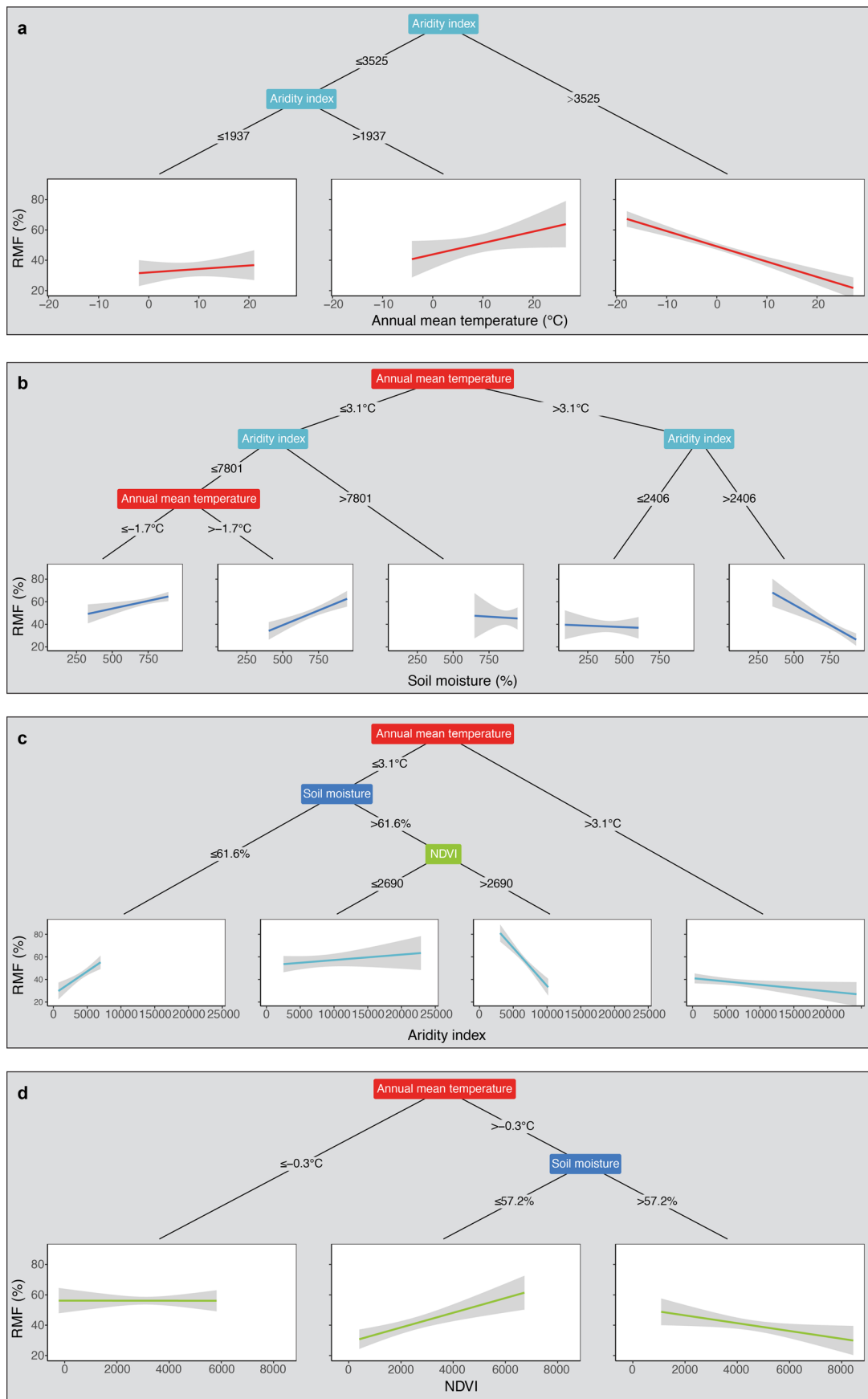
Extended Data Fig. 4 | See next page for caption.

Extended Data Fig. 4 | Recursive partitioning trees for the univariate effects of annual mean temperature (a), soil moisture (b), NDVI (c), and sand content (d) on RMFs in forests. These four variables were chosen on basis of the random forest variable importance metric (Fig. 3a) and, for each model, the remaining three variables were evaluated as potential split points. The number of independent observations contained in each terminal node was constrained to $\geq 10\%$ of the total data (500 observations). Regression plots show slopes and 95% confidence intervals.



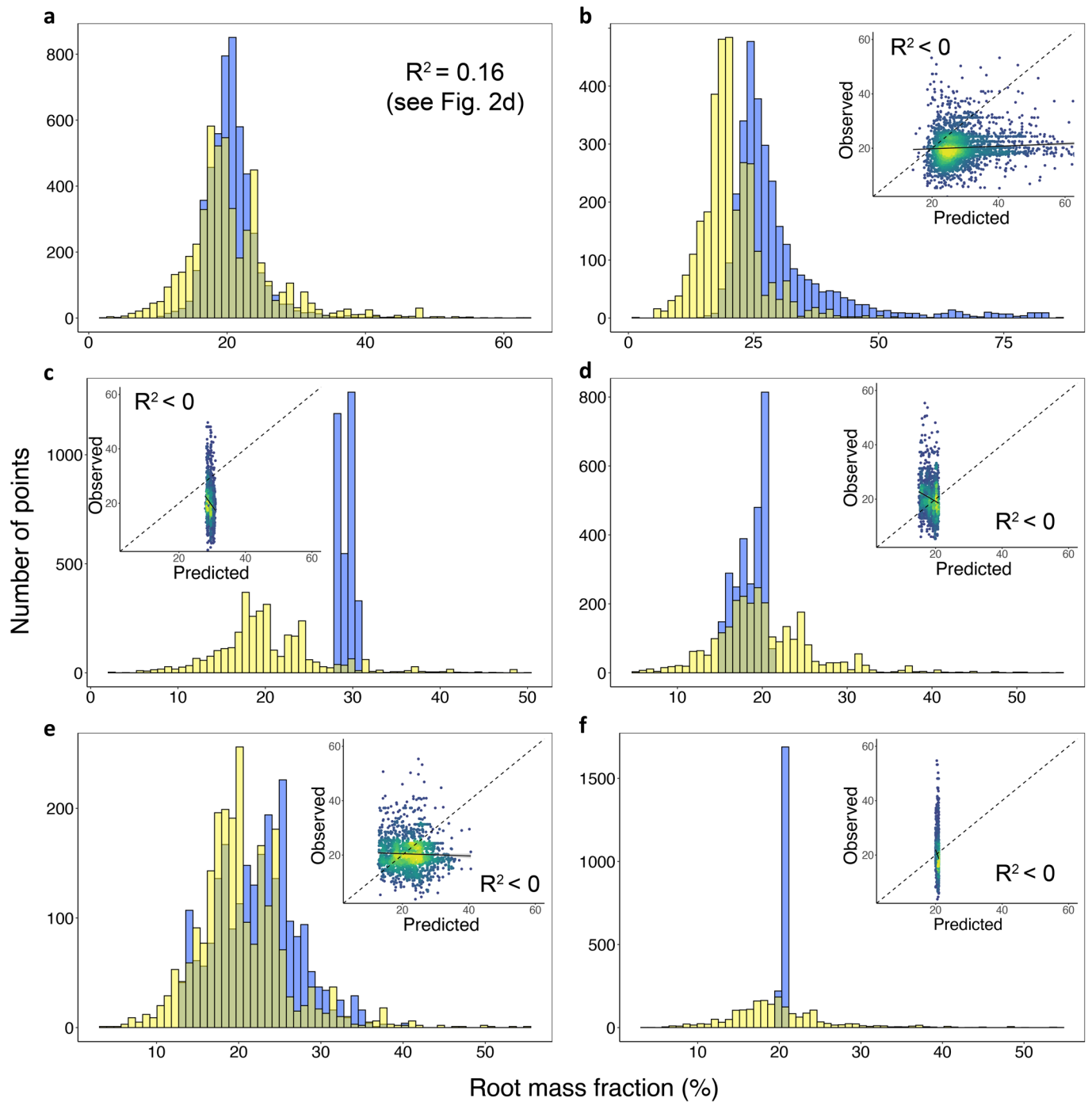
Extended Data Fig. 5 | See next page for caption.

Extended Data Fig. 5 | Recursive partitioning trees for the univariate effects of annual mean temperature (a), soil moisture (b), aridity index (c), and NDVI (d) on RMFs in shrublands. These four variables were chosen on basis of the random forest variable importance metric (Fig. 3b) and, for each model, the remaining three variables were evaluated as potential split points. The number of independent observations contained in each terminal node was constrained to $\geq 10\%$ of the total data (30 observations). Regression plots show slopes and 95% confidence intervals.

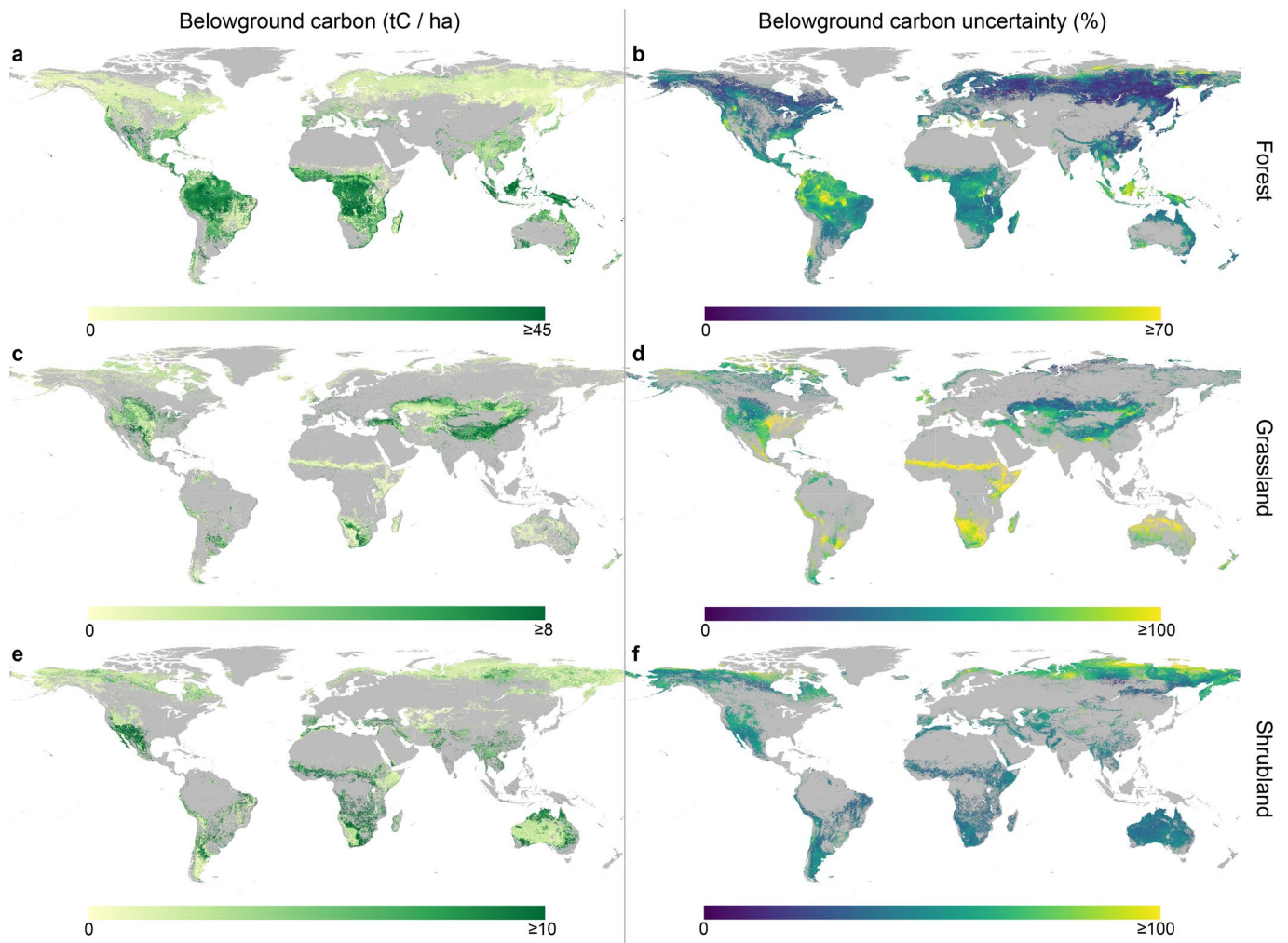


Extended Data Fig. 6 | See next page for caption.

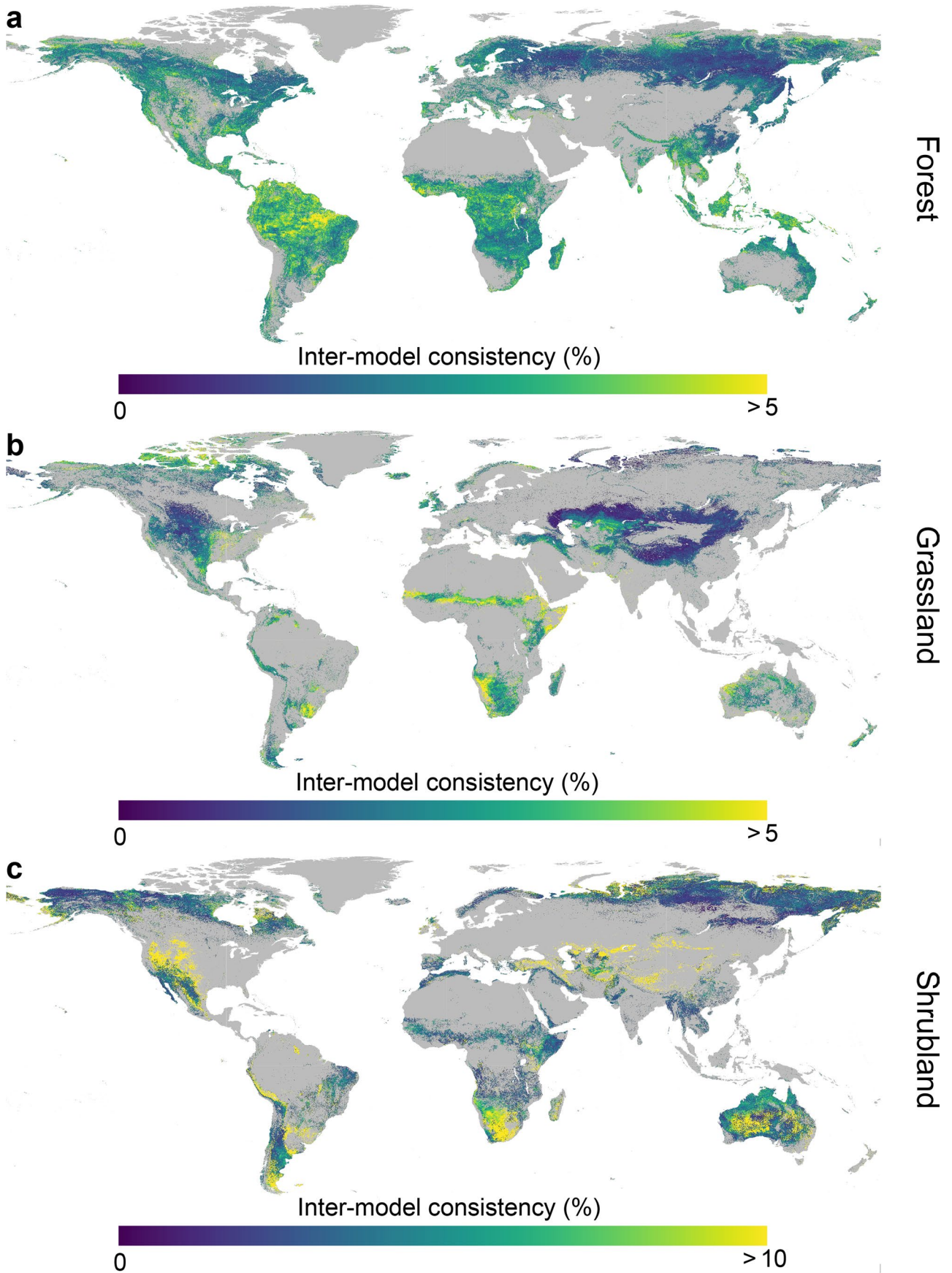
Extended Data Fig. 6 | Recursive partitioning trees for the univariate effects of annual mean temperature (a), soil moisture (b), aridity index (c), and NDVI (d) on RMFs in grasslands. These four variables were chosen on basis of the random forest variable importance metric (Fig. 3c) and, for each model, the remaining three variables were evaluated as potential split points. The number of independent observations contained in each terminal node was constrained to $\geq 10\%$ of the total data (120 observations). Regression plots show slopes and 95% confidence intervals.



Extended Data Fig. 7 | Comparison of observed forest RMFs with predicted RMFs from dynamic global vegetation models and a current-generation biomass map. The blue bars represent histograms of predicted RMF values based on our LOO-CV procedure (**a**), current-generation biomass estimates⁶ (**b**), and the vegetation models CABLE-POP (**c**), CLASS-CTEM (**d**), ISAM (**e**) and ORCHIDEE (**f**). Yellow bars represent observed values. Insets show scatter plots of predicted versus observed RMFs with solid lines indicating fitted relationships, dashed diagonal lines indicating a 1:1 relationship between observed and predicted points. For the vegetation models, forest was defined as pixels with a tree cover fraction higher than 50%.

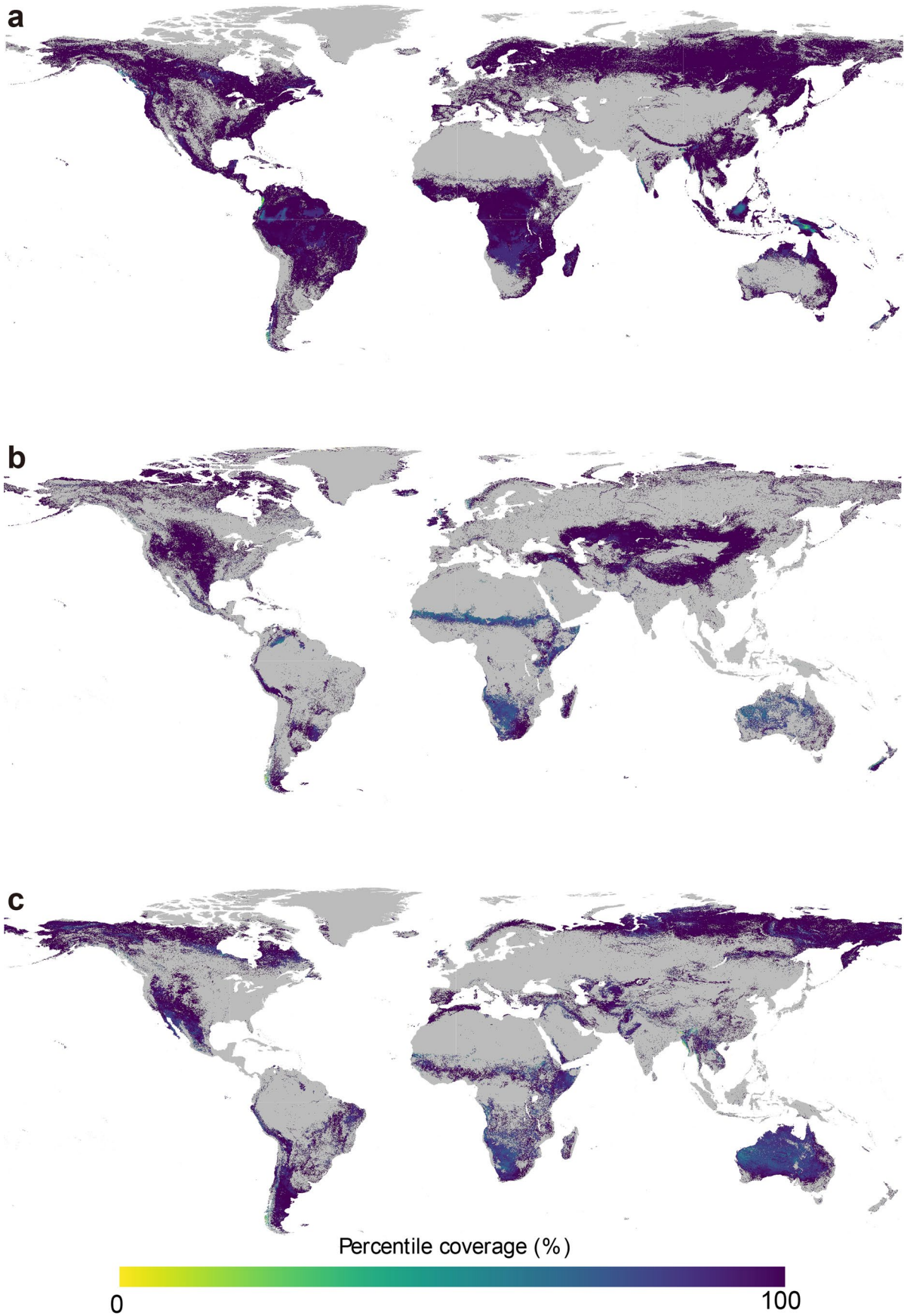


Extended Data Fig. 8 | The global distribution of belowground plant biomass and associated uncertainties in forests (a, b), grasslands (c, d), and shrublands (e, f). a, c, e, Belowground plant biomass (in tons carbon per hectare). b, d, f Associated uncertainties in belowground carbon, calculated as the predicted biomass range (based on 2.5% and 97.5% RMF quantiles derived from the bootstrapped RMF models) divided by the mean predicted biomass in each pixel. Maps are projected at 30 arc-seconds ($\sim 1 \text{ km}^2$) resolution.



Extended Data Fig. 9 | See next page for caption.

Extended Data Fig. 9 | Root mass fraction inter-model consistency in forests (a), grasslands (b) and shrublands (c). Inter-model consistency was calculated as the coefficient of variation (standard deviation divided by mean, in %) of the predictions of the 10 best models. Maps are projected at 30 arc-seconds (~1 km²) resolution.



Extended Data Fig. 10 | See next page for caption.

Extended Data Fig. 10 | The extent of interpolation and extrapolation across all terrestrial pixels in which the respective vegetation type, forest (a), grassland (b) and shrubland (c) occurs. Values represent the percentage of interpolation based on principal component analysis, that is, the percentage of bands that fall into the convex hull space.

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Our web collection on [statistics for biologists](#) contains articles on many of the points above.

Software and code

Policy information about [availability of computer code](#)

Data collection Data was collected manually by searching the literature. Citations for the root-shoot ratio data are provided in the methods.

Data analysis The analyses were run in R [R Core Team (2021) R: a language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. <http://www.R-project.org>. Accessed 17 Jan 2021] and Google Earth Engine (<https://earthengine.google.com>). The code used for this study is available at https://github.com/haozhima95/Global_mapping_root_shoot_ratio.git.

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All manuscripts must include a [data availability statement](#). This statement should provide the following information, where applicable:

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The root-shoot ratio data underlying this study are available at https://github.com/haozhima95/Global_mapping_root_shoot_ratio/tree/master/RSR_data. Citations for the root-shoot ratio data are provided in the methods.

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Ecological, evolutionary & environmental sciences study design

All studies must disclose on these points even when the disclosure is negative.

Study description	We use a global dataset of plant root mass fractions to generate a spatially explicit understanding of the proportional plant biomass that exists in roots and upscale these predictions to the globe to quantify total belowground plant biomass. Our database consists of 17,814 plot-level root mass fraction measurements, aggregated to 6,803 spatially distinct samples from forest (N = 5170, from 52 studies), grassland (N = 1293, from 22 studies), and shrubland (N = 340, from 82 studies) biomes. Using 63 global layers of climate, topography, soil, and vegetation features, we trained random forest models to characterize the environmental factors that drive geographic variations in RMFs within each vegetation type and interpolate these relationships across the globe. We additionally use recursive partitioning and regression methods accounting for spatial autocorrelation and compare our empirical root mass fraction data against vegetation model simulations.
Research sample	Plot-level root-shoot biomass ratios in trees, grasses and shrubs.
Sampling strategy	No statistical methods were used to predetermine sample size.
Data collection	Data was collected through literature search. Citations for the root-shoot ratio data are provided in the methods.
Timing and spatial scale	Root-shoot ratio data from all continents except Antarctica, collected after the year 1960 were used in this study.
Data exclusions	Data was only used if all of the following four criteria were met: 1) RSR values were available at the plot level (and not only for selected individuals); 2) data came from field measurements (allometry-derived RSRs were excluded). If the study consisted of multiple treatments, only data from the control treatment was selected; 3) coordinate information was given; and 4) measurements were taken after 1960. For datasets that included information on more than one vegetation type, the abovementioned criteria were evaluated separately for the subset of trees, grasses and shrubs within the dataset. Furthermore, to minimize the influence of ontogenetic drift on RSRs in forests (higher root-to-shoot biomass ratios in seedlings than in adults ⁴⁵), we 1) excluded studies that only focused on seedlings of trees and 2) excluded seedlings or tree individuals with a DBH smaller than 10cm within studies.
Reproducibility	n/a
Randomization	n/a
Blinding	n/a
Did the study involve field work?	<input type="checkbox"/> Yes <input checked="" type="checkbox"/> No

Reporting for specific materials, systems and methods

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